

STRUCTURAL CONTROLS ON THERMAL ANOMALIES AND DIFFUSE SOIL DEGASSING AT MENENGAI GEOTHERMAL PROSPECT, THE KENYAN RIFT VALLEY

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ABSTRACT

Regional exploration for geothermal resources in Kenya indicates that the Quaternary volcanic complexes of the Kenya rift valley provide the most promising prospects for geothermal exploration. A number of geoscientific studies have been conducted to assess the geothermal potential of this area. This paper reports the results of a survey of CO₂ soil fluxes and concentrations of thoron (Rn₂₂₀) in soil gases. Soil CO₂ concentrations measurements were performed using an Orsat apparatus whereas Rn₂₂₀ soil gas concentrations were measured with a portable radon detector (emanometer). A total of 275 sampling points were measured. The obtained values ranged between zero to 12% of CO₂ in total gas and zero to 6425cpm of Rn₂₂₀.

Peak levels of soil CO₂ concentrations and thoron concentrations were identified to be located on major faults and within the caldera floor where hydrothermal fluids are rising and the rocks are highly fractured allowing the release of CO₂ and thoron locally marked by fumaroles. Areas identified with the highest anomalies for CO₂ and thoron soil gas concentrations can be seen in the contour maps presented. The caldera, Molo TVA and also Solai graben are the important geological structures associated with these anomalies.

Keywords: geothermal resource, thoron concentrations, CO₂ concentrations, TVA and faults

INTRODUCTION

The Menengai Geothermal Prospect is located in the central section of Kenyan Rift Valley, the area north of L. Nakuru and south of Lake Bogoria. The extend of the prospect covers an area of approximately 600 km² characterized by a complex geological setting with Menengai caldera being notably the major geological feature in the area and is also important for its geothermal potential. Most part of this area, in particular the southern section is discussed in this study.

The general exploration for geothermal resources in Kenya indicates that the Quaternary volcanic complexes of the Kenya rift valley provide the most promising prospects for geothermal exploration, and as a result, intensive exploration work has focussed on areas with volcanic centres located within the rift valley. Studies show that these volcanic centres have positive indications of geothermal resource that can be commercially exploited. Menengai geothermal area is one of the priority prospects in the current prospects ranking. Eleven locations with fumaroles have been identified in the region (Figure 1).

A number of thermal anomalies and diffuse degassing assessments have been carried out in most of geothermal prospect areas in Kenya, including Menengai. Results for CO₂ and Rn₂₂₀ soil diffuse concentrations have been found to discharge more in areas that coincide with faults. The purpose of this study is to establish the soil gas concentration of CO₂ and Rn₂₂₀ and relate with known geological structures.

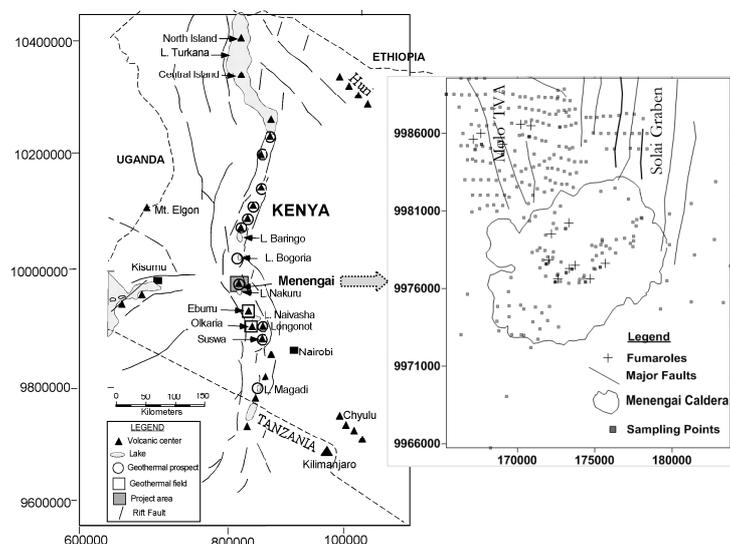


Figure 1. Location of Menengai Geothermal Prospect showing fumaroles, major faults and sampling points.

GEOLOGICAL AND STRUCTURAL PATTERN

The Kenyan Rift is characterized by extension tectonism where the E-W tensional forces resulted in block faulting (KenGen 2004). This included leant blocks as evident in both the floor and scarps of the rift. The rift trough is truncated by several normal faults, which evidently represent persistent and wide-ranging tectonism under the rift floor. Menengai Geothermal Prospect itself is located within an area characterized by a complex tectonic activity associated with the rift triple junction; the latter being a zone at which the failed rift arm of the Nyanza rift joins the main Kenya rift. Two rift floor tectono-volcanic axes (TVA) that are important in controlling the geothermal system in study area include the Molo and the Solai TVA.

Menengai has been active from about 0.4 -0.3 Ma B.P (Gislason 1989). The formation of the shield volcano began about 200,000 years ago and was followed by the eruption of two voluminous ash-flow tuffs, each preceded by major pumice falls. More than 70 post-caldera lava flows cover the caldera floor, the youngest of which may be only a few hundred years old. The caldera of Menengai volcano lies immediately north of Lake Nakuru, but ignimbrites and air-fall tuffs from the volcano cover some 1350km² (Leat & Macdonald 1984) extending into Molo area (Jennings 1971). The volcano is mainly composed of strongly peralkaline, Si-oversaturated trachytes and has had a complex geochemical evolution, resulting from the interplay of magma mixing, crystal fractionation and liquid state differentiation (Leat & Macdonald 1984).

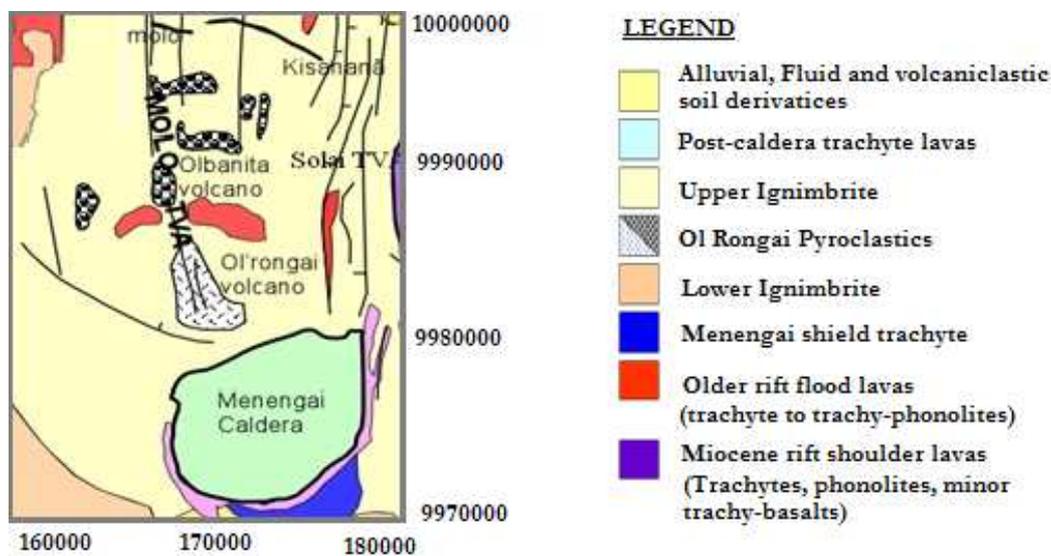


Figure 2. Geological map of Menengai Geothermal Prospect Area (after KenGen 2004)

METHODOLOGY

Early geochemical studies on diffuse soil degassing have been carried out in a number of geothermal areas. CO₂ soil gas investigations similar to the one used in this study have been conducted elsewhere in volcanic areas to examine unexpected rise in CO₂ flux (e.g. Farrar *et al.* 1995; McGee and Gerlach 1998; Gerlach *et al.* 1998), estimate total volcanic flux from volcanic vents and diffuse flank emissions (e.g. Allard *et al.* 1991; Chiodini *et al.* 1996), and to classify tectonic structures related to volcanic degassing (e.g. Giammanco *et al.* 1997; Bergfeld 1998).

The study area together with a total of 275 sampling points is presented in Figure 1. Sampling spacing relied mainly on accessibility of specific targeted points and the ease of getting samples, in particular on the rugged areas. Nonetheless, a random separation of about 300m in average was used. Areas covered by younger lava flow were almost impossible to sample and few samples thus represents such areas, as was seen with lava pond within the caldera. The survey involved measuring of CO₂ and Rn-220 in the soil gas. All the measurements were done at the same point. In addition, soil temperatures were taken using an infra-red thermometer at the point where the soil gas survey was undertaken.

Measurement of CO₂ gas was done by the help of an Orsat apparatus after the soil gas was hand pumped from a depth of 1m below the surface via a hole made using a steel spike with jacket. The Orsat apparatus consist of absorption vessels, which measure about 100 mls and contain a 40% KOH solution for absorbing the acidic CO₂. The corresponding volume changes in the absorption vessel represent the corresponding amounts of the gases in volumes given as percent of the total gas as a percentage.

Radon radioactivity levels were measured in the soil gas using a portable radon detector (emanometer). The soil gas was passed into the emanometer by using a hand-operated vacuum pump and Rn-220 readings recorded in counts per minute (cpm). Three background counts were recorded at three-minute interval prior to introduction of sample into the emanometer. Upon introducing the sample, three more readings were taken at one-minute interval to give the total radon counts.

The CO₂ and Rn₂₂₀ obtained in the survey (both 2004 and 2010 sampling programs) were corrected and used to construct concentrations contour maps (Figures 3 to 6). Natural neighbor contouring method was used because it interpolates scattered and irregularly spaced data effectively and hence produces better contours maps from irregularly spaced measurements (Ledoux & Gold 2005).

RESULTS AND DISCUSSIONS

The results of descriptive statistics of CO₂, Rn₂₂₀ and Temperature collected are presented in Table 1. The frequency distributions are described by normal distribution. The range of values for these parameters spans several orders of magnitude.

Table 1: Descriptive statistics of CO₂, Thoron and Temperature from 275 sampling points.

	N	Range	Minimum	Maximum	Mean	Std. Deviation	Variance	Skewness		Kurtosis	
	<i>Statistic</i>	<i>Error</i>	<i>Statistic</i>	<i>Error</i>							
Rn220	275	6427	0	6427	386	603	363391	5.14	0.15	39.49	0.29
CO2	275	12.30	0.00	12.30	1.12	1.54	2.37	5.12	0.15	30.81	0.29
Temp	275	55.90	21.20	77.10	28.70	6.99	211.52	0.05	0.15	0.62	0.29

Soil Temperature Measurements

The mapping of temperature variations at or below the earth's surface is an essential geothermal exploration instrument. The spatial variation of soil temperature is illustrated in Figure 3. Quite a number of anomalies can be identified in relation to the major structural setting. And as indicated by results, a higher temperature anomaly in the caldera is evident around the fumaroles with the highest fumaroles temperature of 88°C. Similar to other soil surveys, soil temperature results in the caldera is mainly affected by the lava that covers the caldera floor and apart from the areas where the lava did not cover. The areas to the North outside the caldera gave an anomaly confirming the geothermal potential of the area coinciding with two major faults trending NW-SW.

Diffuse Degassing Measurements

The term volcanic gas defines a gas exsolved from a magmatic source of an active volcano whereas hydrothermal gas defines a gas exsolved from the envelope of hot water that surrounds the magmatic environment. Volcanic gases have composition different from the hydrothermal gases, the first is richest in SO₂ and the second in H₂S (Giggenbach 1996). CO₂ occurs in both magmatic and hydrothermal gases and is the most abundant gas after water vapor. The deep-seated faults in the crust tap magmatic CO₂, which is transmitted to the surface where it is naturally lost through the soil. Carbon dioxide (CO₂) of magmatic origin is normally channeled through deep-seated tectonic structures close to the surface of the earth and then seeps out of the ground through the soil. However, dealing with CO₂ anomalies especially in the Rift Valley ought to be treated with caution owing to the availability of localized sources of this gas as suggested by Darling *et al.* (1995). CO₂ may also originate from other sources like organic matter, which are likely to give false impressions of a geothermal source.

Thoron gas (Rn₂₂₀) is a short lived isotope of radon produced by the thorium (Th₂₃₂) decay series with a half-life of 55 seconds (Lopez *et al.* 2004). Due to its short half-life, thoron's transportation is limited to a few centimeters either by diffusive or convective flows, as compared to Rn₂₂₂ (t_{1/2} = 3.8 days) (Hutter 1993). As a result, the high concentrations of thoron are more likely to be due to convective movement of gases rather than diffusive processes and therefore it is difficult to achieve an exact quantitative measurement of its concentration. If we measure the disintegration of thoron, it shows a decrease during the first three minutes after collection of the sample, as opposed to the radon that presents a half-life of almost 4 days (Magana *et al.* 2002). High thoron values could be suggestive of enhanced permeability, as it may as well depend on other factors such as the degree of rock fracture and the ability of groundwater to circulate and remove thoron.

The CO₂ gas concentrations are plotted in Figure 4 while the results of radon emanations were recorded in counts per minute (cpm) for Rn₂₂₀ and plotted in Figure 5. To reduce interferences between CO₂ and Radon sources a ratio of Rn₂₂₀/CO₂ have been plotted in Figure 6.

CO₂ Concentrations

The concentration of CO₂ in the soil gas in the surveyed area is given as a percentage of the total gas in the soil. The degassing of CO₂ through the soil in the prospect area is considered to be mainly supplied by two sources, the one of volcanic origin emanating from deep environment and a shallow one that result from organic activity. In this study, high CO₂ values (>2.5%) are found at a belt running in a NW-SE while other areas exhibit low values (<1 %). These areas coincide with the Molo Tectono-Volcanic Axis (Molo TVA), deep seated faults and the eruption centers in the prospect area which indicate that the source could be from a magmatic body. Highest CO₂ concentration in the soil gas was observed around Ol'Rongai Hill (located along Molo TVA) and the contiguous fumaroles with CO₂>10% which could be indicative of good permeability in the area.

The central part of the caldera around the fumaroles, as well, displays relatively high values of CO₂ (>2.8%). There are also few highs of CO₂>5 % in the western part of the caldera. The presence of thick young lava that covers most parts of the caldera interfere with the movement of CO₂ to the surface and therefore makes sampling in these parts extremely difficult, and is also noticeable in the contouring of the CO₂ distribution as shown in Figure 4. Care ought to be taken while correlating thermal and diffuse degassing anomaly with subsurface permeability. This is because the cap of young lava that covers most part of the caldera might be directing the anomaly laterally beneath the lava bed and hence giving misleading impression as a result.

Thoron Concentrations

From the results of the absolute values of Rn₂₂₀, north-western parts of the surveyed area show the highest concentrations (Figure 5). In addition, a more pronounced anomaly is evident in central part of the caldera in the areas adjoining fumaroles following similar pattern to the one revealed by CO₂ distributions. However, a mild rise in thoron values is seen in the eastern flank touching Solai Graben.

Thoron-CO₂ Ratios

The ratio of the Rn₂₂₀/CO₂ gases would be a good indicator of the magmatic source of the gases since the ratio is not expected to change if they are from the same source (see Figure 6). High values of these ratios are found towards the north, north-western to the north eastern parts outside of the caldera wall with two anomalous areas at the fumaroles in Ol'Rongai area. A mantle source of CO₂ as suggested by Darling et al. (1995) is a possibility of the source of this gas due to intense faulting in the area, and these faults around Ol'Rongai are of great importance in localizing degassing of CO₂ and Rn₂₂₀ as well thermal manifestation anomalies. If a mantle source is inferred due to intense faulting, then the faults should be deep seated. These areas apparently coincide with high CO₂ and high Rn-220 absolute values.

The western part of the surveyed area, which shows high absolute values of CO₂, gives low value of this ratio. This is an indication of a source other than the one of magmatic origin. This coincides with an agriculture farm and consequently the CO₂ anomaly in this area is considered to emanate from organic source, as a result. Nonetheless, a higher ratio of these gases is seen within and to the east of the caldera.

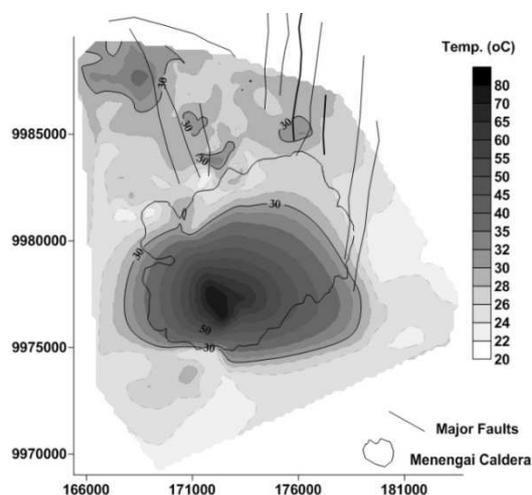


Figure 3: Temperature distribution in °C

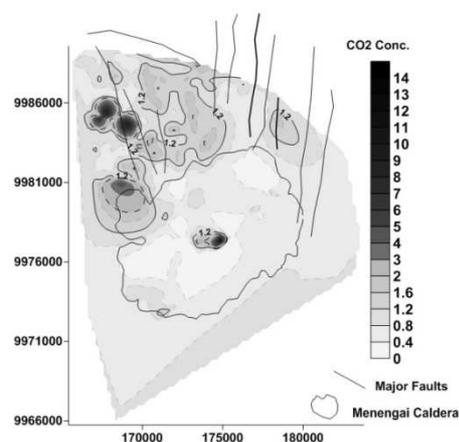


Figure 4: CO₂ conc. distributions

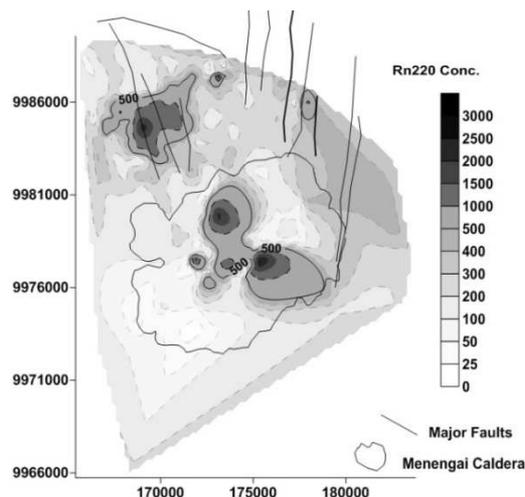


Figure 5: Rn220 conc. distributions

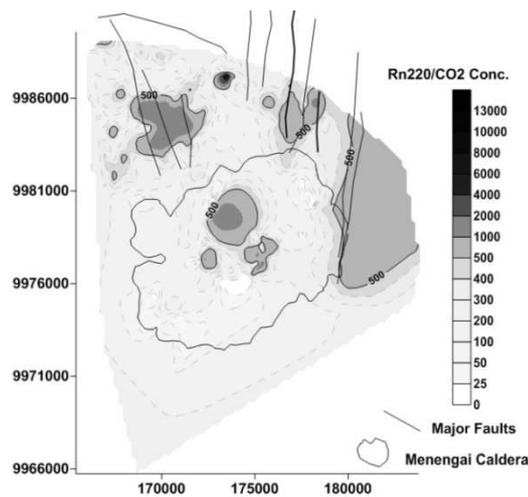


Figure 6: Rn220/CO2 conc. distributions

CONCLUSIONS

Despite the erratic distribution of sampling points, the results from acquired data can be used to make some conclusions relative to the objective. The fumaroles show close relations with geologic structures, which provides means of access for hydrothermal fluids to reach the surface. In the same way, most of the gas anomalies are located in fault zones. CO₂ and Rn₂₂₀ emanations from the presently quiescent faults in Menengai were measured to be higher than background values.

The two fault-lines at Ol'Rongai are considered in this study to be of great importance in geothermal exploration and any future development as it reflects intensity of subsurface permeability in the area, and enhances the confidence to invest. The caldera itself as a geological structure displays a positive evidence of the existence of geothermal resource, with its thermal anomalies that coincides with CO₂ and thoron anomalies. Nonetheless, the picture portrayed by the contours within the caldera is hereby treated with caution as the pond of new lava spread across most part of the floor greatly influenced sampling spaces, and possibility of lateral flow of thermal fluids and diffuse gases at the upper zones restricted by the lava cap cannot be negated.

Generally, CO₂ concentrations are structurally controlled but may have been modified by changes in land use. As seen from the findings, the study of soil CO₂ and Rn₂₂₀ composition as proved to be important geochemical method to identify vertical zones of high permeability. The employment of these methods in other geothermal fields could give important and relatively cheap information for field utilization and development.

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REFERENCES

- Allard P., Carbonnelle J., Dajlevic D., le Bronec J., Morel P., Robe M.C., Maurenas J.M., Faivre-Pierret R., Martin D., Sabroux J.C. & Zettwoog P. 1991. Eruptive and diffuse emissions of CO₂ from Mount Etna. *Nature*, 351, 387–391.
- Bergfeld D., Goff F., Janik C.K. & Johnson S.D. 1998. CO₂ Flux Measurements across Portions of the Dixie Valley Geothermal System, Nevada. *Geothermal Resources Council Transactions*, 22, 20-23.
- Chiodini G., Frondini F. & Raco, B. 1996. Diffuse emission of CO₂ from the Fossa crater, Vulcano Island _Italy. *Bull. Volcanol.*, 58, 41–50.
- Darling W.G., Griesshaber E., Andrews J.N., Armannsson H. and O'Nions R.K. 1995: The origin of hydrothermal and other gases in the Kenyan Rift Valley. *Geochim. Cosmochim. Acta*, 59, 2501-2512.
- Farrar C.D., Sorey M.L., Evans W.C., Howle J.F., Ken B.D., Kennedy B.M., King C.Y. & Southon J.R. 1995. Forest-killing diffuse CO₂ emission at Mammoth Mountain as a sign of magmatic unrest. *Nature* 376, 675-678.
- Gerlach T.M., Doukas M.P., McGee K.A. & Kessler, R., 1998. Three-year decline of magmatic CO₂ emissions from soils of a Mammoth Mountain tree kill: Horseshoe Lake, CA, 1995-1997. *Geophy. Res. Lett.*, 25, 1947-1950.
- Giammanco S., Gurrieri S. & Valenza M. 1997. Soil CO₂ degassing along tectonic structures of Mount Etna (Sicily): the Pernicana fault. *Appl. Geochem.*, 12, 429–436.

- Gislason G. 1989. Terminal Report of Exploration for geothermal energy KEN/82/002, United Nations DTCD.
- Hutter A. R. 1993. Thoron/radon ($^{220}\text{Rn}/^{222}\text{Rn}$) ratios as indicators of soil gas transport: Geological Society of America Abstracts with Programs, p. A 195.
- KenGen 2004. Menengai volcano: investigations for its geothermal potential. A Geothermal Resource Assessment (GRA) project. Internal Report.
- Leat P.T. & Macdonald R. 1984. Geochemical Evolution of the Menengai Caldera Volcano, Kenya. U.S Geology Survey. Journal of Geophysical Research, 89, 8571-8592.
- Ledoux H. & Gold H. 2005. An Efficient Natural Neighbour Interpolation Algorithm for Geoscientific Modelling. In Developments of Spatial Data Handling by Peter F Fisher. 11th International Symposium on Spatial Data Handling, pp. 97-109.
- Lopez D.L., Ransom L., Perez N., Hernandez P. & Monterrosa J. 2004. Dynamics of diffuse degassing at Ilopango Caldera, El Salvador. Rose, W.I., Bommer, J.J., López, D.L., Carr, M.J., and Major, J.J. editors, Geological Society of America Special Paper No. 375 "Natural Hazards in El Salvador, pp.191-202.
- Magaña M., Lopez D., Tenorio J. & Matus A. 2002. Radon and Carbon Dioxide Soil Degassing at Ahuachapan Geothermal Field, El Salvador. Geothermal Resources Council Transactions, 26, 341-344.
- McGee K.A. & Gerlach T.M. 1998. Annual cycle of magmatic CO₂ at Mammoth Mountain, California: Implications for soil acidification. Geology, 26, 463-466.